

Evaluation of Radon (^{222}Rn) Distribution and Its Implications Vis-À-Vis Water Quality of Killiyar River, Kerala, India

G R ANURANI, R LAKSHMI, SABU JOSEPH* and S SUKANYA

Department of Environmental Sciences, University of Kerala, Kerala 695581, India.

Abstract

The discharge of tropical rivers is mainly contributed by the base flow from groundwater especially during summer. Hence, in order to sustain the environmental flow of rivers, the conservation of locations where groundwater discharges into river is a better option than conventional practices viz., redesigning river channel structure and flow regime. Radon (^{222}Rn), a colourless, odourless, inert and natural radioactive noble gas ($t_{1/2} = 3.8$ days), can be used as a proxy to trace the groundwater discharge location/s in the river course. As ^{222}Rn readily dissolves in groundwater, its content in groundwater is relatively higher than surface water. We report here the activity of ^{222}Rn in the river water at ten locations from upstream to downstream of Killiyar river – KR ($n = 6^{\text{th}}$, $L = 24$ km, $A = 102$ km 2), the main tributary of Karamana river, Kerala, India. Surface water samples ($n = 10$) were collected during pre- and post-monsoon of 2017. The radon activity was performed by RAD7, an electronic radon detector (DurrIDGE Company Inc., USA). The activity of radon varied from 157 to 4588 Bq/m 3 in pre-monsoon and 147 to 1740 Bq/m 3 in post-monsoon. The spatial variability of ^{222}Rn activity is observed, and the anomalous high activity location/s indicates groundwater potential in that area. Further, the factors controlling spatial variation of radon were also discussed. Moreover, physico-chemical parameters of river water were also studied. And all the parameters were found to be within the permissible limit of Bureau of Indian Standards (BIS) specifications for potable water (IS -10500: 2012). This is a case study of application of radon for prospecting groundwater potential zones in Killiyar river course, henceforth useful for the water resource management in this riverine environment and is first of its kind in the study region.



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CONTACT Sabu Joseph ✉ jsabu2000@gmail.com 📍 Department of Environmental Sciences, University of Kerala, Kerala 695581, India.



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Introduction

Rivers are the principle source of water supply for larger population in urban and rural parts of Kerala, southwestern State of India. But, many of the small rivers in Kerala get dried during summer and flooded during monsoon seasons. Since water is one of the most vital commodities, the over-exploitation of water resource and water pollution is a problematic condition which ultimately leads to water scarcity problems. Apart from the dissolved or suspended chemical pollutants, the physical factors such as heat and radiations also play considerable impact on aquatic life. River water quality depends on so many factors like catchment area, topography, weather and seasonal variations, degree and nature of developments in the river basin by humans etc.

Exploitation of rivers intended for agricultural, domestic and industrial purposes cause severe adverse effect on rivers of Kerala by means of propagating the organic and inorganic salt content. Groundwater discharge is believed to be dominating for dry season flows in perennial river systems and to sustain aquatic biodiversity.¹ This groundwater seepage is vitally important to the hydrologic settings of the world for eternal flow of water in rivers during the times of no rainfall. The interaction between groundwater and surface water is quite complex process.² It depends on many parameters like geology, geomorphology and climate of the particular area. Nevertheless, it is important to measure the groundwater discharge into rivers for getting better knowledge about groundwater balances and for determining sustainable limits of groundwater extraction thereby protecting the environmental flows of river.

Radon (²²²Rn), a natural tracer that has been used in many hydrological studies^{3,4,5} is produced by the alpha decay of ²²⁶Ra in the decay series of ²³⁸U. ²²²Rn (Z=86) is a radioactive element under noble gas family and chemically inactive.^{3,4,5} ²²²Rn is the most suitable water tracer among the 37 radioactive isotopes of radon (¹⁹³Rn to ²²⁹Rn) due to its relatively longer half- life (3.8 days). It is well known that radon activity concentration in groundwater is 2-4 times higher in magnitude compared to that in surface waters.^{6,7,8} This is due to the fact that groundwater is in contact with mineral grains with ²²⁶Ra and ²³⁸U.^{9,10} Most of the

radon produced within a mineral grain remains embedded in the grain, and 10–50 % escapes to enter the pore space.^{11,12} This is called the emanation coefficient, another important factor contributing to higher activity of radon in groundwater.^{13,14,15} The emanation coefficient depends on the type of rock and its structure and porosity.^{13,14,16} Because of the presence of water in the pore space, the radon atoms can dissolve in groundwater. But in case of surface water, turbulence facilitates rapid gas exchange of radon with atmosphere.¹⁷ These properties enable the use of ²²²Rn for studies regarding interaction between groundwater and surface water.^{3-5,18-21} Kies *et al.*²² suggests that radon migrates through the pores in soils, cracks and fractures in rocks such as faults and thrust along with the groundwater which consequently flows into surface water bodies.

Thus, groundwater discharge points can be explored in locations in a river where there is significant increase of radon activity. Therefore, ²²²Rn can be used for groundwater prospecting in a river course and such studies are sparse in Kerala. Hence, from a hydrological perspective, a holistic approach to understand the environmental flow in Killiyar river (KR), Kerala has been attempted in this study using radon as a groundwater tracer.

Materials and Methods

Study Area

Killi river (KR), also known as Killiyar (n=6th, L=24 km, A=102 km²), the major tributary of Karamana River, is located between latitudes 8°40'30" N, 8°27'0" N and longitudes 76°57'0" E, 76°2'0" E in Thiruvananthapuram district of Kerala State, India (Fig. 1). It has its origin near Panavur (8°38'30.7" N and 76°59'19.4" E) physiographically in the midland (7.5-75m elevation amsl) and the river confluences with Karamana river at Kalladimukham (08°27'23.4" N and 76°57'32" E) in downstream region and about 2.0 km inland from the sea coast. The river flows through an undulating terrain in dendritic to sub-dendritic pattern.

Geologically, 90% of river basin is composed of garnetiferous-biotite-sillimanite gneiss with or without graphite (i.e., khondalite), migmatites and some patches of charnockites, sand and clay deposits.

Methodology

Ten sampling locations (Table 1) were selected (interval= ~ 3 km) from upstream to downstream and the sampling was done on a single day during pre- and post-monsoon seasons (January 2017 and December 2017). The sampling sites were selected based on accessibility, land use pattern, topography, degree of anthropogenic interventions at these locations etc. Water temperature, electrical

conductivity (EC), pH, total dissolved solids (TDS), salinity and resistivity were measured onsite using Horiba LAqua U-50 multi-parameter water quality portable kit.²³ Other parameters viz., alkalinity, chloride, total hardness, nitrate, phosphate, sulphate, sodium and potassium were analysed through laboratory experiments using standard procedures of APHA.²⁴

Table 1: Portrait of Sampling Stations, Killiyar

Station ID	Sampling Station	Latitude	Longitude
S1	Puthenpalam	08°37'39.6" N	77°00'02.3" E
S2	Pazhakutty	08°36'46.4" N	76°59'48.1" E
S3	10thstone Bridge	08°35'18.4" N	76°59'43.0" E
S4	Karakulam	08°34'13.9" N	76°59'16.2" E
S5	Vazhayila	08°32'50.3" N	76°58'29.1" E
S6	Mannamoola	08°31'59.4" N	76°58'52.5" E
S7	Maruthankuzhi	08°30'46.3" N	76°58'39.2" E
S8	Jagathy	08°29'29.1" N	76°57'54.4" E
S9	Attukal	08°28'25.0" N	76°57'27.5" E
S10	Kalladimugham	08°27'23.4" N	76°57'32.0" E

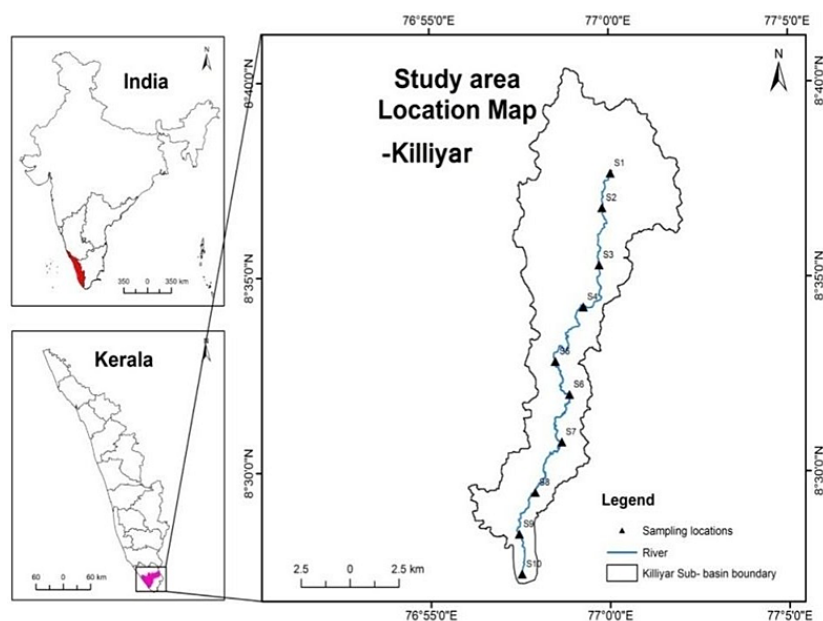


Fig.1: Location map of Killiyar river basin

The surface water samples for radon (^{222}Rn) measurement were collected carefully in 250 mL glass bottles without atmospheric air contact. In laboratory, ^{222}Rn dissolved in water is stripped

by bubbling air and circulated through a closed air-loop via a desiccant tube into the ^{222}Rn counting system (RAD7, Durrige make). The air continuously recirculates through the water and extracts more

than 95% of the dissolved radon until a state of equilibrium develops i.e., within 5 minutes. The radon monitor consists of setup with a high electric field above a silicon semiconductor detector at ground potential to attract the positively charged polonium daughters, $^{218}\text{Po}^+$ and $^{214}\text{Po}^+$, which are counted as a measure of ^{222}Rn activity expressed in Bq/m^3 (disintegration per second per m^3). The measured radon activities were corrected for radioactive decay since water sampling. The correlation of ^{222}Rn with physico-chemical characteristics of water and correlation among different parameters were analysed using SPSS.¹⁷

Results and Discussion

Radon Activity Distribution

The spatial variation of radon (^{222}Rn) activity from upstream (S1; 0 km) to downstream (S10; 24 km)

along the course of Killiyar river (KR) during pre-monsoon (PRM) and post-monsoon (POM) is given in Table 2 and Figure 3. The spatial distribution of radon followed almost similar pattern during both seasons. During PRM, ^{222}Rn activity varied from 157 to 4588 Bq/m^3 and during POM, it ranged from 147 to 1740 Bq/m^3 . The results demonstrated spatial variation of radon in which the upstream reaches of KR (S1-S2; 0-3 km stretch) exhibited reasonable higher radon activity ($>1000 \text{ Bq}/\text{m}^3$) in contrast to the remaining downstream reaches of river (except S5) with lower radon activity ($<500 \text{ Bq}/\text{m}^3$) in both seasons (Fig. 3). However, the location S5 in the mid course of KR ($\sim 10 \text{ km}$ downstream of S1) also exhibited higher ^{222}Rn activity ($>1000 \text{ Bq}/\text{m}^3$) for both seasons (Fig.3).

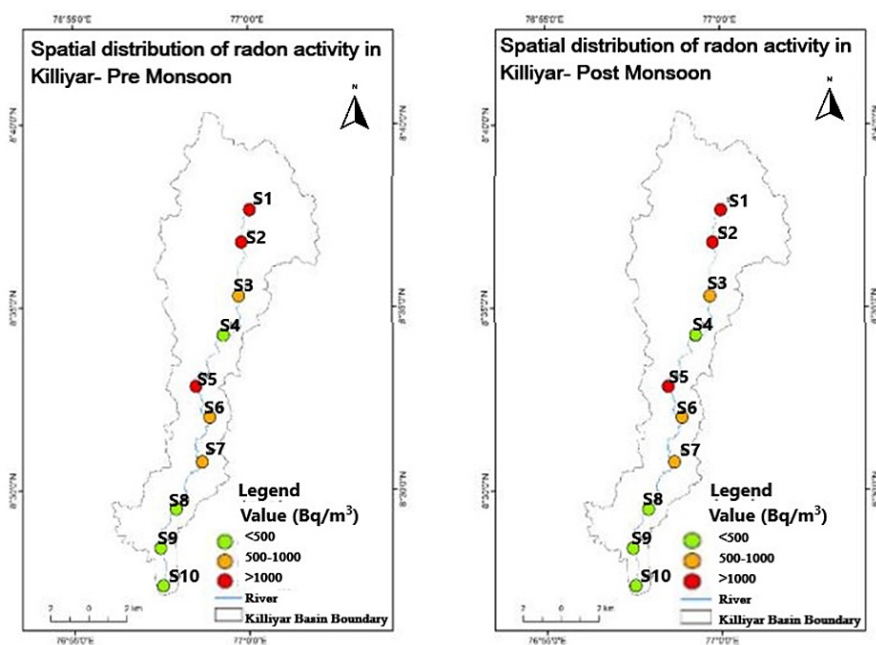


Fig. 2: Map showing spatio temporal variation of radon activity in Killiyar river during Pre-monsoon and Post-monsoon of 2017

Generally, increased ^{222}Rn activities in surface water are indicative of radon-rich groundwater influx into the river.²⁵ Hence, the upstream (S1-S2; 0-3 km stretch) and midstream (S5; 10 km downstream of S1) stretch in the river, where anomalous high radon activities ($>1000 \text{ Bq}/\text{m}^3$) noticed, are likely zones of prospective groundwater discharge into the river. The ^{222}Rn activity shows a decreasing

trend from upstream to downstream (except at S5), suggesting the escape of radon from surface water to atmosphere across the boundary layer through molecular diffusion. i.e., gas exchange.²⁶ Another possible reason for this decline towards downstream is radioactive decay of ^{222}Rn due to short half-life ($t_{1/2} = 3.8 \text{ days}$).²⁷

Further, temporal variation of radon activities do exist in the study area. i.e., the activity of radon found to be lower during POM (mean=745 Bq/m³) when compared to PRM (mean=1150 Bq/m³). The dilution effect of monsoon is likely to have influenced in the decrease in activity of radon in river during POM. Additionally, the turbulent condition of water after monsoon favours fast atmospheric escape of radon.

Factors Influencing Radon Activity In Water

Major factors governing the spatial variation of surface water radon along the river were scrutinized.

Geological Control on Radon Variability

The source of radon is geogenic²⁸⁻³⁰ and its content is primarily controlled by uranium and radium present in the host rock. ²²²Rn activities vary spatially based on geology of the region (i.e., bed rock types),

degree of weathering, radon releasing potential (i.e., emanation coefficient), structural features in rocks (i.e., presence of fractures), lineaments in the study area etc.^{15, 28,29} Moreover, near surface weathering and bedrock fracturing are implicated with increased radon activities.³⁰

Interestingly, locations with higher ²²²Rn activities in the river course (i.e., S1-S2, 0-3 km stretch and S5- 10 km from S1) are underlain by fractured compact khondalitic rocks susceptible to weathering. These rock formations do not possess primary porosity, and tend to bear groundwater in repositories with the development of secondary porosity due to weathering and fracturing and by other tectonic processes.³¹ The distinctiveness of these rocks is that, the quantity of groundwater available is primarily dependent on the storage and rate of infiltration through the fractures and further relies on the nature of fractures. The foliations of these metamorphic rocks serve as planes of weakness and facilitate flow and storage of groundwater.³²

Table 2: Radon activity in Killiyar during pre-monsoon and post-monsoon

Sampling Stations	Radon Activity in Bq/m ³	
	Pre-monsoon	Post-monsoon
S1	1788	1470
S2	4588	1740
S3	684	633
S4	265	204
S5	1373	1180
S6	943	736
S7	938	599
S8	468	442
S9	157	294
S10	312	147

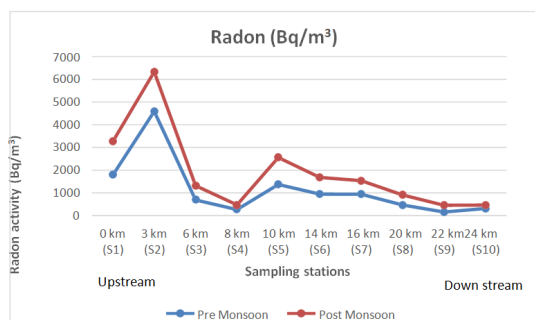


Fig. 3: Graphical representation of spatio-temporal variation of ²²²Rn activities in Killiyar

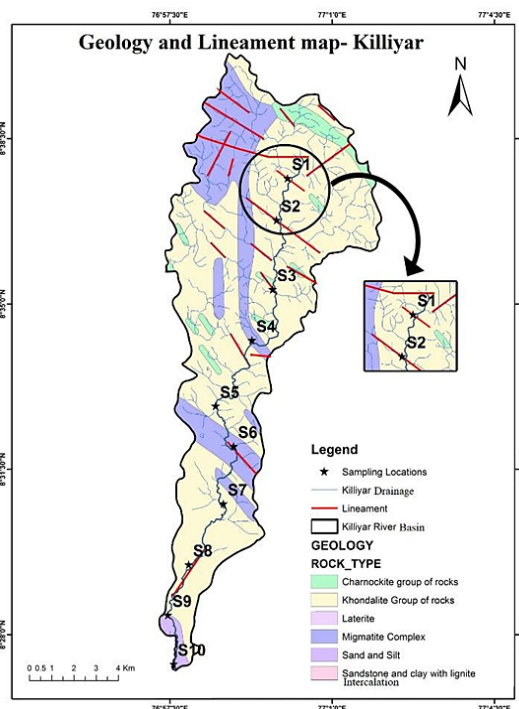


Fig. 4: Map showing major geology and lineaments in Killiyar River Basin along with sampling locations

However, migmatitic rocks were also noticed in few river reaches (e.g., S4 - 8 km downward of S1; S6- 14 km downward of S1) with relatively lower ^{222}Rn activities. i.e., 265-943 Bq/m³ during PRM and 204 -736 Bq/m³ during POM.

Further, river course underlain by sand and silt in downstream (S9 & S10; 22-24 km stretch) was found to have the lowest ^{222}Rn activities with 157 to 312 Bq/m³ during PRM and 147-294 Bq/m³ during POM. This may be attributed to low content of radon parent minerals (uranium bearing minerals) in sandy silt alluvium. Further, the river bed is silty clay underlain by thick sandy silt layers. This silty clay unit is having low hydraulic conductivity compared to the underlying sand and silt layer.³³ A similar condition was reported in Murray-Darling river basin, Australia by Lamontagne *et al.*³⁴ This distinction of fluvial geology and associated ^{222}Rn activities indisputably explains the role of geology in subsurface input of groundwater.

Structural Influence on Radon Activity

Apart from geology, gradient of the region and lineaments associated with tectonic disturbances also show variable role in contributing the inflow of radon to surface water.³⁵ Lineaments are regions where potential groundwater development possible and it provides secondary porosity to the rocks and serve as conduits for rapid movement of radon from aquifer to the river. Regions where lineaments intersect each other are expected to be more favourable for groundwater development.^{36,37}

Lineament map (Fig. 4) of KR river basin shows two major sets of lineaments dominating the region. One set is NW-SE trending and the other is NE-SW trending. Interestingly, the NW- SE trending lineament transect the river near two sampling sites in the upstream (S1 and S2; 0-3 km stretch), where ^{222}Rn activities are high and hence, this active lineament present here could be one of the reasons for lineament-controlled radon and groundwater discharge in those regions (0-3 km stretch).

The river stretch at 10th km from upstream (S5) is devoid of lineaments, but show high radon activity and indicates factor/s other than lineament take a role. However, previous studies by Sreeja *et al.*³³ developed a groundwater potential zone map of KR basin based on geospatial data integration of

lineament layer, geology, geomorphology, drainage and slope of the terrain, and indicated that the S5 location is a potential groundwater development zone. Thus, a combination of factors contributes to high radon at S5.

Activity of radon tends to be low in the downstream part of KR, even though some locations (e.g., S3 and S8) are underlain by khondalite group of rocks. The associated lineaments (NE-SW) in this area might be passive lineaments, and further the less weathered and fractured rocks might also be the reason for low radon.

Groundwater Potential Zones Based on Radon Activity

Based on the radon activity, the groundwater potential zones of KR are classified into three categories - high, moderate and low or negligible groundwater potential zones (Fig. 2). This is based on the interpretive guide put forth by Harrington *et al.* 2012.³⁸ The high radon activity (>1000 Bq/m³) zones are at stretches 0-3 km upstream (S1 and S2) and 10th km from upstream (S5) and implies high groundwater prospects. The medium radon activities (500- 1000 Bq/m³) recorded at 6th km from upstream (S3) and between 14th and 16th km from upstream (S6 & S7), implies moderate groundwater potential zone. The low radon activities (100-500 Bq/m³) implying poor groundwater potential zone is spread in the upper reach 8th km from upstream (S4) and in the lower segment of river from 20-24 km (S8-S10).

Geophysical (Resistivity) field surveys can be applied in this zone of KR river course for groundwater prospecting and confining the groundwater potential in these zones. Once it is confined, then further management options can be charted out to preserve the groundwater potential in these zones for the sustainable environmental flow of the river during dry season.

Surface Water Quality

The measured values of environmental variables of surface water samples are presented in Table S1.

Most of the parameters varied spatio-temporally. All water quality parameters analyzed at locations upstream to downstream were within BIS permissible limits.³⁹

The pH ranged from 6.1 to 7.4 during PRM (mean = 6.87 ± 0.41) and 5.1 to 6.8 during POM (mean = 6.06 ± 0.42). The surface water samples were acidic in nature during POM suggesting increased free CO_2 from acidic rainwater contribution to the river during monsoon.

Highest EC values were observed at S10 (24th km location in downstream) consistently during both seasons (PRM=3550 & POM=1229 $\mu\text{S}/\text{cm}$) indicating anthropogenic pollution as well as salt water intrusion to lower reaches through the estuary.

Total alkalinity in the river water ranged between 60 and 180 mg/L during PRM (mean = 90 ± 34.3) and from 88 to 226 mg/L during POM (mean = 157 ± 43.6). Elevated levels of total hardness were witnessed in the lower most reaches during both seasons (PRM =344; POM =165 mg/L) contributed by saline water encroachment and active washing ghats.

The concentrations of chloride increased remarkably above BIS acceptable limit 250 mg/L (BIS, 2012)^{26, 35} at S10 during both seasons (Table.3) implying the anthropogenic inputs and estuarine influence viz., seawater intrusion, sea sprays, tidal effects.

Nitrate levels ranged from 1.94 to 15.68 mg/L (mean = 7.3 ± 4.5) in PRM and 2.83 to 19.40 mg/L (mean = 7.7 ± 6.4) in POM. Nitrate concentration showed an analogous trend during both seasons with relatively increased levels at three locations i.e., S2, S7 and S10. It is attributed to the discharge of domestic wastes which acts as a favorable substrate for biological oxidation of nitrogenous organic matter thus by releasing nitrate. Also, nutrient loaded in Karamana river joins with Killiyar river at S10 location in the downstream.⁴⁰ Dilution due to heavy rains during monsoon and related runoff account for the declined concentration of PO_4 during POM (mean = 0.47 ± 0.13) in all stations compared to PRM (mean = 4.02 ± 2.6).

Sulphate concentrations ranged from 3.49 to 6.32 mg/L in PRM and 2.58 to 8.8 mg/L in POM. The S10 location showed comparatively high content of EC, TDS, salinity, alkalinity, chloride, hardness, NO_3 , PO_4 and Na mainly due to anthropogenic pollution and salt water intrusion.

Correlation of Radon with Physico-Chemical Parameters

In general, the correlation analysis of various water quality parameters with radon showed no significant correlation for both seasons (Table S1, S2). Similar observation was noted by Idriss *et al.*⁴¹ It endorses the fact that radon in water is attributed to physical processes and geological conditions of a region rather than chemical interactions.^{15, 42, 43}

Effect of Temperature

Temperature is one of the critical factors governing ^{222}Rn dissolution in water⁴⁴ and its solubility increases with decreasing temperature.⁴⁵ However, this dissolution rate is detectable only when there is considerable temperature differences i.e., 30-50°C between the samples.^{15, 46} Radon showed no correlation with water temperature in KR. This could be explained by the insignificant temperature difference within the river water samples (26-29°C).

Effect of pH

The results of the correlation analysis suggested that there is no relationship between pH and radon. Becker⁴⁷ observed a positive correlation between ^{222}Rn and pH in samples with $\text{pH} > 8$, which is higher than the pH obtained in Killiyar surface water samples (range = 5.3-7.4). Further, Ye *et al.*⁴⁸ stated that under experimental conditions, the mean water radon solubility ($K_{w/air}$) is as low as 0.24 when pH increases from 3 to 13 thus by confirming that there is no significant effect of pH on radon dissolution.

Effect of Electrical Conductivity

In Killiyar river, Electrical Conductivity (EC) showed no correlation with radon. Similar results were observed in previous studies.^{15, 49} However, few other studies^{50, 51} observed a weak positive correlation of radon with EC and moderate negative correlation with temperature. Conversely, Akawwi⁵² observed a negative correlation of radon with EC.

Correlation among Water Quality Parameters

The correlation matrix (Table S1 and Table S2) shows relationship between different parameters in both seasons. EC is highly negatively correlated with resistivity in both seasons ($r = -0.95$ in PRM and -0.92 in POM). TDS showed high positive correlation with Na in both season ($r = 0.96$, $p < 0.01$) and with K ($r = 0.98$ in PRM and 0.90 in POM). Salinity is

positively correlated with Na ($r=0.97$ in PRM and 0.93 in POM) and K ($r=0.97$ in PRM and 0.90 in POM, $p < 0.01$). Cl also showed very high positive correlation with Na in both season ($r=0.97$, $p < 0.01$) and with K ($r=0.96$ in PRM and 0.90 in POM, $p < 0.01$). Na showed high positive correlation K ($r=0.97$ in PRM and 0.91 in POM, $p < 0.01$). The correlation among the foresaid parameters (EC, Na, Cl, K) reflect a combined influence of marine inputs (seawater intrusion, marine sprays, tidal effects), anthropogenic inputs and agricultural runoff into the hydrochemistry of KR.

The discharge of tropical rivers is mainly contributed by the baseflow from groundwater during summer. The minimum flow that a river should have in order to preserve its ecosystems and environmental quality is known as 'environmental flow (e-flow)'.⁵³ In Killiyar, one major challenge is the difficulty to sustain this e-flow during summer, and hence e-flow should be rejuvenated through an integrated holistic river management approach with community participation. The conservation of groundwater discharge sites on the river banks is one viable option. Further, communities on the river bank should be encouraged for aquifer recharging through various rainwater harvesting techniques. Another challenge is quality of riverine water towards downstream due to rapid urbanization. Unscientific waste dumping from human settlements and indiscriminate discharge of sewages directly into the river, reflects the local community's inadequate exposure to proper waste management practices. Proper waste management-cum-awareness programme is a viable option for this.

Conclusion

This study characterizes ²²²Rn activity and different water quality parameters in Killiyar river (KR), Kerala (India) from upstream to downstream during two seasons, viz., pre- monsoon (PRM) and post-monsoon (POM). The radon activity showed spatial variations from upstream to downstream with anomalous higher values in upstream locations (S1, S2) and midstream (S5), and are likely zones of potential groundwater discharge into the river. Seasonal variation showed high radon activity during

pre-monsoon (mean= 1150 Bq/m^3) compared to post-monsoon (745 Bq/m^3) and is due to dilution from monsoon. Based on the study, three groundwater potential zones were identified in the river, viz., high (²²²Rn $> 1000 \text{ Bq/m}^3$), medium (500-1000) and low ($100\text{-}500 \text{ Bq/m}^3$) potential zones.

From the water quality analysis, it is found that, most of the parameters in most locations (except S9 and S10) were within permissible limits of BIS. In S10 most of parameters are relatively high due to anthropogenic contribution and also marine influence through estuary. Further, radon activity is not correlated with any of water quality parameters in the study area.

From the study, it is concluded that ²²²Rn isotopic application was used as a case study in Killiyar river and groundwater potential zones in Killiyar river course were identified. Methods like geophysical (resistivity) field surveys can be applied in this zone of river course for further groundwater prospecting and confining the groundwater potential. Further, stable isotopes (¹⁸O, ²H) and supplementary data viz., resistivity, piezometric levels, river discharge etc. could be applied to quantify groundwater discharge into river using mass balance approach. This baseline information would benefit local authorities and planners for implementing integrated holistic river management programmes for the sustainable environment flow of river especially during pre-monsoon season with community participation.

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Conflict of Interest

The authors do not have any conflict of interest.

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